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1 **Remote sensing of photosynthetic light-use efficiency across two forested biomes: Spatial scaling**

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49 **Abstract**

50 Eddy covariance (EC) measurements have greatly advanced our knowledge of carbon exchange in  
51 terrestrial ecosystems. However, appropriate techniques are required to upscale these spatially discrete  
52 findings globally. Satellite remote sensing provides unique opportunities in this respect, but remote  
53 sensing of the photosynthetic light use efficiency ( $\epsilon$ ), one of the key components of Gross Primary  
54 Production, is challenging. Some progress has been made in recent years using the photochemical  
55 reflectance index, a narrow waveband index centered at 531 and 570nm. The high sensitivity of this  
56 index to various extraneous effects such as canopy structure, and the view observer geometry has so far  
57 prevented its use at landscape and global scales. One critical aspect of upscaling PRI is the development  
58 of generic algorithms to account for structural differences in vegetation. Building on previous work, this  
59 study compares the differences in the PRI: $\epsilon$  relationship between a coastal Douglas-fir forest located on  
60 Vancouver Island, British Columbia, and a mature Aspen stand located in central Saskatchewan, Canada.  
61 Using continuous, tower-based observations acquired from an automated multi-angular  
62 spectroradiometer (AMSPEC II) installed each site, we demonstrate that PRI can be used to measure  $\epsilon$   
63 throughout the vegetation season at the DF-49 stand ( $r^2=0.91$ ,  $p<0.00$ ) as well as the deciduous site  
64 ( $r^2=0.88$ ,  $p<0.00$ ). It is further shown that this PRI signal can be also observed from space at both sites  
65 using daily observations from the Moderate Resolution Imaging Spectro-radiometer (MODIS) and a  
66 multi-angular implementation of atmospheric correction (MAIAC) ( $r^2=0.54$  DF-49;  $r^2=0.63$  SOA;  $p<0.00$ ).  
67 By implementing a simple hillshade model derived from airborne light detection and ranging (LiDAR) to  
68 approximate canopy shadow fractions ( $\alpha_s$ ), it further demonstrated that the differences observed in the  
69 relationship between PRI and  $\epsilon$  at DF-49 and SOA can be attributed largely to differences in  $\alpha_s$ . The  
70 findings of this study suggest that algorithm used to separate physiological from extraneous effects in  
71 PRI reflectance may be more broadly applicable and portable across these two climatically and  
72 structurally different biome types, when the differences in canopy structure are known.

73 **1. Introduction**

74 Global and spatially continuous estimates of plant photosynthesis are required for a comprehensive  
75 understanding of the terrestrial carbon cycle and the determination of CO<sub>2</sub> uptake by plants (Barr et al.  
76 2004). Over the last few decades, eddy covariance measurements of CO<sub>2</sub> exchange between the canopy  
77 surface and its surrounding air column have greatly improved our understanding of carbon cycling at the  
78 stand level (Baldocchi 2003); however, appropriate techniques are required to upscale these findings to  
79 landscape and global scales (Chen et al. 2003; Reichstein et al. 2007). Satellite remote sensing offers  
80 unique opportunities in this respect, through provision of a globally continuous parameterization of the  
81 land surface at regular time intervals from space (Hall et al. 2005).

82 Gross primary production (GPP) of green vegetation is proportional to the photosynthetically active  
83 radiation (PAR [MJ]) incident upon the canopy at a given time, the fraction of it being absorbed by the  
84 green vegetation elements ( $f_{PAR}$ ) and the efficiency  $\epsilon$  [g CMJ<sup>-1</sup>] with which plants can use this absorbed  
85 radiation energy to produce biomass (Monteith 1972, 1977). This efficiency, also known as light-use  
86 efficiency, is driven by any of a large number of factors restraining the photochemical reaction process,  
87 such as temperature, nutrient and water supply and, as a result, varies greatly in space and time (Field  
88 and Mooney 1986). One of the most common methods used for remote sensing of  $\epsilon$  is the  
89 photochemical reflectance index (PRI) (Gamon et al. 1993; Gamon 1992), a narrow waveband,  
90 normalized difference index that relates  $\epsilon$  to a xanthophyll-induced absorption feature at 531 nm, which  
91 is intimately linked to the biochemical mechanism down-regulating photosynthesis and dissipating  
92 excessive radiation energy as heat (Demmig- Adams and Adams 1996). While the relationship between  
93 PRI and  $\epsilon$  has been proven across a wide range of species (Filella et al. 1996; Gamon et al. 1993;  
94 Garbulsky et al. 2008; Penuelas et al. 1995), its generalization to satellite observable scales is  
95 challenging, as PRI is also driven by numerous other factors including the sun-observer geometry, soil  
96 background reflectance, canopy structure and the ratio of carotenoid to chlorophyll concentration( also

97 referred to as pigment pool size) (Asner 1998; Barton and North 2001; Hall et al. 2008; Hilker et al.  
98 2008a). In addition to the uncertainties existing at the close range, spaceborne observations of PRI are  
99 also confounded by atmospheric scattering (Drolet et al. 2005; Drolet et al. 2008; Hilker et al. 2009b).  
100 Although these effects can generally be accounted for by modeling the radiative transfer of light  
101 through the atmosphere (Vermote and Kotchenova 2008; Vermote et al. 1997), the simplifying  
102 assumptions underlying the commonly used, single orbit-based atmospheric correction algorithms,  
103 cause uncertainties in the PRI wavebands (Hilker et al. 2009b), whose total change in reflectance  
104 between relaxed and photo-inhibited state is in the order of only about 6% (Hall et al. 2008).

105 Using a tower-mounted, automated multi-angular spectro-radiometer (AMSPEC), Hilker et al. (2008)  
106 introduced a technique to separate the extraneous effects from the physiological signal contained in  
107 stand level PRI which allowed, for the first time, a temporally continuous remote sensing of  $\varepsilon$ . Year-  
108 round reflectance data were stratified into observations taken under homogenous physiological and  
109 atmospheric conditions and the bi-directional reflectance distribution function (BRDF) was determined  
110 separately for each stratum. It was then shown that the physiological component of the canopy-level PRI  
111 signal was contained in the change of BRDF adjusted reflectance across strata (Hilker et al. 2008a) which  
112 were directly linked to changes in the xanthophyll cycle of vegetation (Hall et al. 2008).

113 At the satellite scale, Drolet et al. (2005) introduced a first spaceborne assessment of  $\varepsilon$ , using data  
114 acquired from the Moderate Resolution Imaging Spectroradiometer (MODIS). A relationship between  
115 the normalized difference of MODIS bands 11 and 12 ( $PRI_{12}$ ) and EC-measured-  $\varepsilon$  was found when  
116 restricting data to backscatter observations (Drolet et al. 2005). Similar studies since confirmed these  
117 findings (Drolet et al. 2008; Goerner et al. 2009). Building on the work of Drolet et al. (2005, 2008),  
118 Hilker et al. (2009b) used AMSPEC data to “translate” EC-measured  $\varepsilon$  into a stand-level PRI signal first,  
119 which was then compared to MODIS observations after adjusting the viewing geometries of the two  
120 sensors. A new, multi-angular implementation of atmospheric correction (MAIAC) algorithm (Lyapustin

121 and Wang, 2009) was used to correct for atmospheric scattering which, for the first time, allowed the  
122 use of forward and backward scatter observations. Previously, the atmospheric noise in the MODIS  
123 standard product, and an incomplete correction for BRDF effects masked the weaker forward scatter PRI  
124 changes with LUE variations. The MAIAC-corrected MODIS PRI markedly enhanced the relationship  
125 between MODIS and tower-based observations throughout the year (Hilker et al. 2009b).

126 One critical aspect for the development of a more generic algorithm that allows remote sensing of  $\varepsilon$   
127 across the landscape and eventually at global scales, is the study of species and structure related  
128 differences in PRI (Gamon et al. 1993). For instance, Barton and North (2001) found that PRI is sensitive  
129 to species related differences in leaf angle distribution and leaf area. Similarly, (Gamon et al. 1997)  
130 found statistically significant differences in the mean annual PRI across a range of different plant  
131 functional types. (Sims and Gamon 2002) found PRI observations to be sensitive to variations in the  
132 pigment pool sizes existing across species and over time. In this study, we assess and compare the  
133 differences in the relationship between PRI and  $\varepsilon$  across two forested biomes using data simultaneously  
134 acquired at the Douglas-fir (*Pseudotsuga menziesii var menziesii* (Mirb.) Franco) stand and a mature  
135 Aspen forest located in Prince Albert National Park, Saskatchewan, Canada. First, we demonstrate that  
136 the approach previously used to establish a year round, stand-level relationship between PRI and  $\varepsilon$  at  
137 the Douglas-fir site (hereafter DF-49) (Hilker et al. 2008a) can successfully be applied also at the Aspen  
138 stand. Second, tower-borne PRI data acquired at both sites are related to spaceborne observations  
139 taken from the MODIS sensor (Hilker et al. 2009b) and the relationships are compared between the two  
140 sites. Finally, the differences between the PRI: $\varepsilon$  relationships observed at the coniferous and deciduous  
141 stand are being investigated and quantified using a LiDAR derived model of the canopy surface to assess  
142 mutual shading effects of individual tree crowns. Our results show that the differences in the PRI: $\varepsilon$   
143 relationship between the sites can be attributed to differences in canopy shadow fractions ( $\alpha_s$ ) existing  
144 at the coniferous and deciduous stand. Furthermore, we show that the instantaneous derivative of PRI

145 with respect to  $\alpha_s$ , computed from multi-angle sensor measurement, can be used to infer instantaneous  
146 canopy  $\epsilon$  for both the Douglas-fir and Aspen sites using a single functional relation. Our results suggest  
147 that multi-angle remote sensing techniques of  $\epsilon$  might be species invariant thus applicable across  
148 vegetated landscapes without detailed knowledge of vegetation structure and composition.

149

## 150 **2. Methods**

### 151 **2.1 Study areas**

152 The DF-49 site is a 61-year old, second-growth coniferous forest located on Vancouver Island, British  
153 Columbia, Canada, at 300 m above sea level (49°52'7.8" N, 125°20'6.3" W). The stand consists of 80%  
154 Douglas fir, 17% western red cedar (*Thuja plicata* Donn ex D. Don) and 3% western hemlock (*Tsuga*  
155 *heterophylla* (Raf.) Sarg.) and is among the most productive in Canada (Morgenstern et al. 2004). The  
156 stand density is 1100 stems ha<sup>-1</sup>, with tree height ranging between 30 and 35 m. The site is located  
157 within the dry maritime Coastal Western Hemlock bio-geoclimatic subzone (mean annual temperature  
158  $\approx 8.5^\circ\text{C}$ ), which is characterized by cool summers and mild winters with occasional drought during late  
159 summer (Humphreys et al. 2006). The leaf area index (LAI) is 7.3 m m<sup>-2</sup> (Chen et al. 2006).

160 The mature Aspen study site, hereafter referred to as Southern Old Aspen (SOA), has been established  
161 as part of the Boreal Ecosystem-Atmosphere Study (BOREAS) carried out between 1994 and 1996 and is  
162 located in central Saskatchewan (53.62889° N, 106.19779° W, altitude 600 m). The 86-year old stand is  
163 situated in the southern ecotone of the western boreal forests (mean annual temperature  $\approx 0.5^\circ\text{C}$ ) and  
164 consists of trembling aspen (*Populus tremuloides* Michx) with about 10% of balsam poplar (*Populus*  
165 *balsamifera* L.) and a thick, 2-3 m hazelnut understory (*Corylus cornuta* Marsh) with sparse alder (*Alnus*  
166 *crispa* (Alt.) Porsch) (Barr et al. 2007). A 1998 stand survey found the stem density was 830 stems ha<sup>-1</sup>,  
167 the mean tree height of the overstorey is about 22 m (Barr et al. 2007), and the mean LAI is 2.1 m m<sup>-2</sup>  
168 (Chen et al. 2006).

169 **2.1 Eddy covariance measurements**

170 Continuous, half-hourly fluxes of CO<sub>2</sub> have been acquired at DF-49 and SOA as part of the Canadian  
171 Carbon Program (Margolis et al. 2006). Net ecosystem exchange (NEE) was determined as the sum of  
172 the half-hourly fluxes of CO<sub>2</sub> and the rate of change in CO<sub>2</sub> storage in the air column between ground  
173 and EC measurement level, using a three-axis sonic anemometer-thermometer (Model R3, Gill  
174 Instruments Ltd., Lymington, UK, both sites) and a closed-path CO<sub>2</sub>/H<sub>2</sub>O infrared gas analyzer (LI-6262,  
175 LI-COR Inc., Lincoln, NE, USA, both sites) (Barr et al. 2004; Jassal et al. 2007). Incident and reflected PAR  
176 [ $\mu\text{mol m}^{-2} \text{s}^{-1}$ ] was measured from upward and downward looking quantum sensors (model 190 SZ and  
177 190SA, LI-COR Inc. at DF-49 and SOA, respectively) above and below the canopy and  $f_{PAR}$  was derived at  
178 each site from the incident and reflected total PAR measured above and below the canopy, leaf area  
179 index, and the solar zenith angle ( $\theta$ ) at the time of measurement (Chen 1996; Chen et al. 2006). Gross  
180 primary production (GPP) was determined as the difference between NEE and daytime ecosystem  
181 respiration ( $R_D$ )(Humphreys et al. 2006) and  $\epsilon$  was calculated as (Monteith 1972, 1977)

182 
$$\epsilon = \frac{GPP}{PAR \times f_{PAR}} \quad (1)$$

183

184 **2.2 LiDAR data acquisition**

185 Discrete return airborne LiDAR data were acquired at the DF-49 site on August 14th 2008, using a Leica  
186 ALS50-II recording up to 4 returns per outbound laser pulse. The sensor pulse rate was 110 kHz, at an  
187 approximate flying altitude of 900 m. The estimated GPS accuracy of the sensor was 0.02, 0.03 and 0.05  
188 m in x, y and z, respectively. When both ground and non-ground returns were considered, the dataset  
189 had an average density of 3.74 pts m<sup>-2</sup>. Ground and non-ground returns were separated using a series of  
190 algorithms appropriate for the ground topography (Kraus and Pfeifer 1999) and canopy height model



191 was generated at a spatial resolution of 1 m (Fusion v 2.65, USDA, Forest Service). See (Coops et al.  
192 2007) for more details.

193 A second multiple return ( $\leq 4$ ) airborne LiDAR data collection was acquired by the Applied Geomatics  
194 Research Group, Nova Scotia at the OA site on August 3rd, 2008 using an Optech Inc. (Toronto, Canada)  
195 Airborne Laser Terrain Mapper (ALTM) 3100. The survey was configured using a pulse repetition  
196 frequency (PRF) of 71 kHz, a flying altitude ranging between 700 m and 800 m, and a scan angle of  $\pm 20$   
197 degrees from nadir. A 50% flight line swath overlap was used, resulting in a point density of  
198 approximately 10 returns per  $m^2$ . All multiple return point positions were post-processed relative to a  
199 nearby GPS base station located over a survey monument within 30 km of the survey area. Following  
200 integration of sensor position, attitude and laser range data, the point cloud data were tiled, outlying  
201 points were filtered, a bundle-adjustment or strip-matching procedure was applied to all flight lines, and  
202 the 'cleaned' point-cloud was classified into "ground", "non-ground", and "all" returns using TerraScan  
203 software (TerraSolid, Finland). Validation flights performed over a previously surveyed airport runway  
204 prior to and following the data collection demonstrated that RMS errors in point cloud elevations were  
205 within 10 cm. A digital elevation model (DEM) was created from the ground classified returns using a  
206 triangulated irregular network (TIN) interpolation procedure. This surface was then subtracted from a  
207 digital surface model (DSM) of the all hits returns, which was generated using an inverse distance  
208 weighted (IDW) interpolation procedure. The resultant difference surface was a canopy height model  
209 (CHM) at a resolution of 1 m grid cell spacing.

210

## 211 **2.3 Tower-based spectral observations**

### 212 *2.3.1 AMSPEC II system*

213 Canopy spectra were automatically obtained at both sites using AMSPEC II (Hilker et al. submitted) an  
214 enhanced version of AMSPEC (Hilker et al. 2007). The instrument now features a pan-tilt unit which  
215 allows the sensor head to be moved at any zenith angle ( $\theta_v$ ) between  $43^\circ$  and  $78^\circ$  (view azimuth ( $\phi_v$ )  
216 between  $0$  and  $360^\circ$ , Figure 1). To allow sampling under varying sky conditions, canopy spectra were  
217 obtained from simultaneous measurements of solar irradiance and radiance, sampled every 5 seconds  
218 from sunrise to sunset at a  $10^\circ$  angular step width (horizontally and vertically), thereby completing a full  
219 rotation every 15 minutes. The spectro-radiometer used is a Unispec-DC (PP Systems, Amesbury, MA,  
220 USA) featuring 256 contiguous bands with a nominal band spacing of 3nm (full width half maximum 10  
221 nm) and a nominal range of operation between 350 and 1200 nm. The upward pointing probe is  
222 equipped with a cosine receptor (PP-Systems) to correct sky irradiance measurements for varying solar  
223 altitudes. AMSPEC II also allows tracking of satellite orbits (Crawford et al. 1996; Kelso 2007), thereby  
224 driving the radiometer probe to mimic the satellite viewing geometry during each overpass at the site of  
225 installation (Hilker et al. submitted). While the probe movements are limited by the physical boundaries  
226 of the pan-tilt unit especially for higher satellite elevations, the feature can help stabilizing the BRDF  
227 models used to match the viewing geometries of satellite and tower-based measurements (Hilker et al.  
228 submitted).

229 Two identical units were built and installed at DF-49 and SOA, respectively. The DF-49 system was  
230 installed on May 14, 2009 at a height of 42m ( $\approx 10$  m above the tree canopy) on an open-lattice type 0.5  
231 m triangular flux-tower. No observations were made for  $190^\circ \leq \phi_v \leq 280^\circ$  due to obstruction by the  
232 tower. AMSPEC observations at DF-49 include AMSPEC II data sampled between May 14<sup>th</sup> and October  
233 20<sup>th</sup> 2009 and older, AMSPEC I data (same radiometer but at a fixed zenith angle of  $\theta_v = 63^\circ$ ) sampled  
234 between April 1<sup>st</sup> 2006-March 31<sup>st</sup> 2007, and March 17 – October 21<sup>st</sup>, 2008 (Hilker et al. 2007).

235 The SOA system was installed on May 26, 2009 at a height of 37 m ( $\approx 15$  m above the tree canopy) on a  
236 2.9 m double-scaffold tower. The range of azimuth angles obstructed by the SOA tower was  $90^\circ \leq \phi_v \leq$   
237  $180^\circ$ . At SOA, AMSPEC II data were sampled between May 26 and November 4<sup>th</sup>, 2009.

238

### 239 *2.3.2 Determining seasonality from phenological camera data*

240 A fundamental difference between the two sites is the seasonal change in phenology at the coniferous  
241 and deciduous stand. While the evergreen DF-49 stand is driven by a temperate climate, with tree  
242 growth occurring throughout year (Morgenstern et al. 2004), the deciduous stand is subject to distinct  
243 seasonality and the growing season is determined by spring green-up and leaf senescence in fall. The  
244 phenological state of deciduous canopies exerts a major control on spatial and temporal patterns of GPP  
245 (Richardson et al. 2007), and as a result, seasonal changes in the canopy were expected to greatly affect  
246 the spectral observations sampled at SOA (Kodani et al. 2002). Additionally, reflectance observations in  
247 the spring and late fall were expected to be strongly affected by soil background effects the reflectance  
248 of non-photosynthetically active parts of the canopy. In this study we focus on the spatial aspects of  
249 scaling PRI, SOA observations were restricted to the relatively stable growth period during summer,  
250 while seasonal and temporal changes will be discussed in a second, forthcoming study.

251 One of the improvements implemented in AMSPEC II is a webcam system that is installed in parallel to  
252 the downward pointing probe and automatically samples an image with every spectrum that is observed  
253 by the radiometer (Figure 1) (Hilker et al. submitted). This system was used to track the phenological  
254 changes in the plant canopy by quantifying the divergence of the red and the blue channel from the  
255 brightness observed in green channel of the camera (Richardson et al. 2007):

$$256 \quad 2G_{RB_i} = 2\mu_G - (\mu_R + \mu_B) \quad (2)$$

257 where  $\mu_G$ ,  $\mu_R$  and  $\mu_B$  are the camera observed brightness values (raw DN) in the green, red and blue  
 258 channel, respectively. Richardson et al. (2007) introduced an approach to define the seasons in  
 259 deciduous vegetation by fitting a sigmoid function to the brightness values observed in  $2G_{RB_i}$  and  
 260 using its inflection points to mark the beginning and end of the season. In this study, we adapted the  
 261 method of Richardson et al. (2007) to the slightly more complex patterns found at SOA which are  
 262 determined by an earlier green-up of the Hazelnut understorey and a secondary green-up of the Aspen  
 263 overstorey (Barr et al. 2004; Griffis et al. 2004). As a result, a 4<sup>th</sup> order polynomial rather than a sigmoid  
 264 was selected to fit the  $2G_{RB_i}$  observations throughout the observation period and null and inflection  
 265 points were determined using its first, second and third derivative.

266 As the webcam observations are also affected by directional and sun illumination effects, one  
 267 observation was extracted per day (around solar noon) at a fixed viewing direction to minimize the BRDF  
 268 effects on the camera data.  $\phi_v$  was selected to observe an intermediate amount of shadow within the  
 269 canopy ( $\phi_v = 65^\circ$ ) while  $\theta_v$  was set to a off-nadir direction to minimize potential background  
 270 reflectance effects ( $\theta_v = 73^\circ$ ) (Richardson et al. 2007).

271

### 272 *2.3.2 Separating directional and physiological effects on PRI reflectance*

273 The physiological signal contained in multi-angular, canopy-level PRI observations can be separated from  
 274 extraneous effects when stratifying data into homogenous conditions with respect to the physiological  
 275 and atmospheric conditions under which they were observed (Hilker et al. 2008a). Within each stratum,  
 276 the BRDF of PRI can then be modelled as the linear combination of isotropic, geometric and volumetric  
 277 scattering components (Hilker et al. 2008a; Roujean et al. 1992):

$$278 \quad PRI(\theta_v, \theta_s, \Delta\phi) = k_i + k_g F_g(\theta_v, \theta_s, \Delta\phi) + k_v F_v(\theta_v, \theta_s, \Delta\phi) \quad (3)$$

279 where  $\theta_s$  and  $\Delta\phi$  are the view zenith and relative azimuth angle between sun and observer, respectively;  
280  $k_i$ ,  $k_g$  and  $k_v$  are the isotropic, geometric and volumetric scattering coefficients, and  $F_g$  and  $F_v$  represent  
281 the geometric and volumetric scattering kernel functions, respectively.

282 The physiological status of the canopy was determined at SOA and DF-49 using EC-measured  $\varepsilon$  and the  
283 atmospheric conditions were assessed by modelling the clear-sky solar irradiance as a function of  $\theta_s$  and  
284 comparing it to the irradiance measured by AMSPEC at a given time (Hilker et al. 2009a). Observations  
285 were stratified in steps of  $0.1 \text{ gCMJ}^{-1}$  and 10<sup>th</sup> percentiles of potential sky irradiance, respectively (Hilker  
286 et al. 2008a). Geometric and volumetric scattering were modelled at both sites using the Li-Sparse (LS)  
287 and Ross-Thick (RT) kernels based on a geometric-optical approach of (Li and Strahler 1985) and the  
288 radiative transfer theory of (Ross 1981).

289

#### 290 **2.4 MODIS data acquisition and atmospheric correction**

291 Daily level 1B (L1B) at-sensor radiances (Collection 5) on board the EOS-Aqua and Terra spacecrafts were  
292 acquired for the DF-49 and SOA from the Land Processes Distributed Active Archive Center (LPDAAC)  
293 (data portal: <https://lpdaac.usgs.gov>) for all clear days during the study period and atmospherically  
294 corrected using MAIAC (Hilker et al. 2009b; Lyapustin and Wang 2009). The MAIAC algorithm is based on  
295 multi-orbit retrievals of calibrated top-of-atmosphere reflectance to simultaneously retrieve  
296 atmospheric and surface reflectance parameters, such as aerosol optical thickness (AOT), spectral  
297 regression coefficient (SRC) and spectral surface BRDF (Lyapustin and Wang 2005). The time series  
298 approach of MAIAC, which directly retrieves surface BRF from measurements, has been shown to yield  
299 significantly enhanced relationships between spaceborne and tower-measured PRI as compared to  
300 conventional atmospheric correction approach based on a single-orbit data and Lambertian assumption  
301 (Hilker et al. 2009b).

302 MODIS observes the land surface under different viewing geometries, and consequently, the spatial  
303 extent of the pixels, or “footprint” varies with each overpass. In order to simplify the handling of MODIS  
304 observations, MODIS data are routinely “gridded” to 1 x 1km raster based on a forward and inverse  
305 mapping approach which includes the spatially weighted reflectance of adjacent MODIS pixels (Wolfe et  
306 al. 1998). While this process greatly simplifies data handling, it also introduces uncertainties to the  
307 surface reflectance as the spatial origin of a reflectance measurement becomes less well defined (Tan et  
308 al. 2006). In order to assess these uncertainties on PRI reflectance at the two sites, two types of MODIS  
309 observations were processed and compared in this study, the gridded 1 km standard product and non-  
310 gridded (swath) data (Hilker et al. 2009b).

311

#### 312 *2.4.1 Adjusting the viewing geometries of MODIS and AMSPEC*

313 One advantage of using the tower-measured PRI observations rather than comparing EC-measurements  
314 to MODIS spectra directly, is the possibility to adjust the differences in viewing geometry between the  
315 two sensors (Hilker et al. 2009b). Retrieval of accurate BRDF for PRI wavebands from MODIS is difficult,  
316 as multiple orbits are required to obtain a sufficient number of different sun-observer geometries,  
317 during which the canopy reflectance may change as a result of xanthophyll induced changes in PRI<sub>12</sub>  
318 (Hilker et al. 2008a). AMSPEC completes a full sweep of the forest canopy every 15 minutes. During this  
319 time period, the physiological status of the canopy is assumed to be constant (Hilker et al. 2009b). Half-  
320 hour observations ( $\pm 15$  minutes from peak elevation of the satellite) were extracted from AMSPEC data  
321 during each MODIS overpass at the SOA and DF-49 site and a separate BRDF was modelled for each  
322 overpass using the Roujean approach (Eqn. 3) (Hilker et al. 2009b).

323 MODIS features a band centered at 531 nm (Band 11) which is sensitive to xanthophyll detection, but  
324 lacks a suitable reference band at 570 nm (Gamon et al. 1992). This reference band may, however, be

325 substituted using MODIS band 12, a narrow reflectance band centered at 551 nm (Drolet et al. 2005;  
326 Drolet et al. 2008; Hilker et al. 2009b). The MODIS-based PRI ( $PRI_{12}$ ) is defined as (Drolet et al. 2005)

$$327 \quad PRI_{12} = \frac{\rho_{11} - \rho_{12}}{\rho_{11} + \rho_{12}} \quad (4)$$

328 where  $\rho_{11}$  and  $\rho_{12}$  is the reflectance of MODIS band 11 and 12, respectively. In order to make AMSPEC  
329 observations more comparable to MODIS, AMSPEC derived spectra were resampled to simulate the 10  
330 nm resolution of the MODIS bands 11 and 12 using the arithmetic mean of the corresponding  
331 spectroradiometer wavelengths and  $PRI_{12}$  observations were derived also from AMSPEC data.

332

## 333 **2.2 Estimation of canopy shading**

334 Under conditions where photosynthesis is limited by factors other than light, sunlit parts of the canopy  
335 are exposed to more excessive radiation energy than those shaded by other vegetation elements. Hall et  
336 al (2008) showed that under these conditions, canopy level PRI is strongly dependent on  $\alpha_s$ , and that the  
337 directional changes observed in PRI at a given half hour interval can be attributed almost entirely to  
338 changes in  $\alpha_s$  (Hall et al. 2008). The same study also showed that the slope of the relationship between  
339  $\alpha_s$  and PRI ( $\Delta\alpha_s \Delta PRI^{-1}$ ) changes as a function of  $\epsilon$  and that PRI shows no variation with  $\alpha_s$  when  
340 photosynthesis is not down-regulated (Hall et al. 2008). As a result, the instantaneous derivative of PRI  
341 with respect to  $\alpha_s$  can be used to infer canopy light-use efficiency. The rate of change in  $\Delta\alpha_s \Delta PRI^{-1}$   
342 should be invariant to species related differences between PRI and  $\epsilon$  because Hall et al. (2008) showed  
343 theoretically and empirically that  $\Delta\alpha_s \Delta PRI^{-1}$  is invariant to non-photosynthetically active canopy  
344 elements. These elements, however, are a major driver of spectral differences observed between  
345 species.

346 One simple way to approximate  $\alpha_s$  at least under clear sky conditions is using a hillshade algorithm (Hais  
347 and Kucera 2009) based on a CHM such as available from LiDAR. While the method takes into account  
348 only the mutual shading of tree crowns, Hilker et (2008b) has shown it may still be used to derive  
349 realistic estimates of  $\alpha_s$  incident upon a canopy at a given time. First, the portions of the canopy visible  
350 to AMSPEC were determined by means a viewshed (Kim et al. 2004) applied to the LiDAR derived CHM  
351 at SOA and DF-49. Second, a hillshade was applied to model illumination conditions of the visible parts  
352 of the canopy areas based upon slope, exposition derived from the CHM and  $\theta_s$  and  $\phi_s$  at the time of  
353 observation. The instantaneous field of view of AMSPEC was approximated as an ellipse given by  $\theta_v$  and  
354  $\phi_v$  and the height of installation above canopy (h). For each AMSPEC observation,  $\alpha_s$  was determined as:

$$355 \quad \alpha_s = 1 - \frac{\sum_1^n \kappa}{n_{(\theta_v, \phi_v)}} \quad (5)$$

356 where  $\kappa$  is the modelled brightness of a visible pixel in the hillshade raster (scaled between 0 and 1) and  
357  $n_{(\theta_v, \phi_v)}$  is the total number of visible pixels contained in the field of view of AMSPEC at a given time.

358 One limitation of this LiDAR derived assessment of  $\alpha_s$  is that it can only be applied under clear sky  
359 conditions (Hilker et al. 2008b) as the model does not account for diffuse sky radiation. In order to  
360 assess species related changes between PRI and  $\varepsilon$ , AMSPEC data were extracted from the two clearest  
361 days of each month (as determined by the sum of total daily PAR measured at each site) and used to  
362 determine  $\Delta \text{PRI} \Delta \alpha_s^{-1}$  for each 15 minute interval of these days.

363

### 364 **3. Results**

365 Figure 2 shows daily estimates of vegetation green-up and leaf-down observed by AMSPEC's webcam  
366 system during the 2009 study period. The seasonal dynamics in the  $2G_{RB}_i$  were much stronger at the  
367 Old Aspen site (Figure 2A), compared to the DF-49 site, where almost no changes canopy greenness were  
368 observed (Please note that gap in Figure 2B is due to an instrument downtime at DF-49 between DOY



369 197 and DOY 231). The 4<sup>th</sup> order polynomial function selected to quantify the seasonal changes at SOA  
370 fitted the camera observations well ( $r^2=0.72$ ,  $p<0.01$ ). The minimum camera measured  $2G_{RB_i}$  at this  
371 site was observed at around DOY 175. Up until then, the camera data showed a decreasing trend. After  
372 DOY 175, the webcam observed a substantial green-up of the canopy, which peaked at around DOY 280.  
373 Using null and inflection points of the polynomial function shown in Figure 2A, analysis of AMSPEC  
374 observations at SOA was restricted to DOY 175 - 308. Given the little variation in canopy greenness  
375 observed at DF-49, all available spectra were used at this site.

376 The relationship between EC-measured  $\varepsilon$  and AMSPEC observed, BRDF adjusted PRI and  $PRI_{12}$  is given in  
377 Figure 3. Figure 3A shows the  $PRI:\varepsilon$  correlation observed at SOA (DOY 175 - 308), the corresponding  
378 observations made at the DF-49 site (all 3 years) are presented in Figure 3B. PRI for the sunlit and  
379 shaded part of the canopy is shown (daily averages). At both sites, a highly significant, non-linear  
380 relationship existed between AMSPEC measured PRI and  $\varepsilon$  ( $r^2=0.88$  and  $r^2=0.91$  for SOA and DF-49,  
381 respectively (sunlit canopy),  $p<0.00$ ). At SOA,  $\varepsilon$  measurements ranged between 0 and  $1.8 \text{ gCMJ}^{-1}$  while  
382 PRI measurements, after being adjusted to a common sun-observer geometry, ranged between  
383  $-0.07 \leq PRI \leq -0.02$ . At the same time, the maximum  $\varepsilon$ -value observed at DF-49 was  $2.5 \text{ gCMJ}^{-1}$   
384 while the spectral measurements varied between  $-0.08 \leq PRI \leq -0.02$ . The mean coefficient of  
385 determination for the BRDF models acquired across all strata was  $r^2=0.79$  and  $r^2=0.73$  ( $p<0.00$ ) for SOA  
386 and DF-49, respectively; the standard deviation in both cases was  $\sigma=0.15$ . Figure 3C and D show the  
387 correlation between AMSPEC observed  $PRI_{12}$  and EC-measured  $\varepsilon$  at SOA and DF-49, respectively.  $PRI_{12}$   
388 exposed a similarly significant correlation to  $\varepsilon$  than PRI (Figure 3A and B), the data range, however, was  
389 smaller (Figure C and D) and differences between sunlit and shaded parts of the canopy were less  
390 prominent.

391 The results for upscaling tower-based  $PRI_{12}$  observation to satellite levels are presented in Figures 4 to 6.  
392 Figure 4 shows a BRDF model established from AMSPEC derived  $PRI_{12}$  reflectance during one MODIS

393 overpass (spacecraft noon  $\pm$  15 minutes) as an example. The model presented in Figure 4A shows  
394 observations made at SOA, Figure 4B shows data acquired at DF-49. The x and y-axis in each figure  
395 represent the planar coordinates (origin=tower) of the AMSPEC observations (computed from  $\theta_v$  and  $\phi_v$   
396 and h), the z-axis shows the corresponding  $\rho PRI_{12}$  value. The black dots represent the actual  $PRI_{12}$   
397 measurements of the canopy (for this example:  $n=203$  at SOA and  $n=184$  at DF-49), while the black lines  
398 show the residuals to the fitted BRDF-surface. Overall, the semi-empirical reflectance models described  
399 the directional changes in tower measured  $PRI_{12}$  during the MODIS overpasses well. The average  
400 coefficient of determination was  $r^2=0.93$ ,  $\sigma=0.03$  (SOA) and  $r^2=0.98$ ,  $\sigma=0.05$  (DF49) ( $p<0.00$ ). The red  
401 dots in Figure 4A and 4B represent the  $PRI_{12}$  observations taken by AMSPEC in “satellite tracking mode”  
402 (here tracking the flight path of EOS-TERRA, both figures). The yellow dot in Figure 4A (blue dot in Figure  
403 B) represents the corresponding zenith and azimuth angle of the related MODIS observation. The  
404 different colors of the fitted reflectance surface were used to illustrate the shape of the BRDF model.

405 Figure 5 shows a comparison of the AMSPEC derived, BRDF adjusted MODIS-like PRI based on bands 11  
406 and 12 and PRI from the actual MODIS observations of the same wavelengths (non-gridded  
407 observations, Aqua and Terra spacecrafts combined). The measurements taken at SOA are presented in  
408 Figure 5A, Figure 5B illustrates the reflectance observed at the DF-49 site. While some differences were  
409 found in the absolute reflectance measured by MODIS and AMSPEC (Figure 5A), a significant relationship  
410 between MODIS band 11 and 12 and AMSPEC observed band 11 and 12 existed at both research sites  
411 ( $r^2=0.57$  (SOA, Band 11),  $r^2=0.61$  (SOA, Band 12),  $r^2=0.58$  (DF-49, Band 11),  $r^2=0.62$  (DF-49, Band 12);  
412  $p<0.01$ ). A strong correlation also existed between the  $PRI_{12}$  measurements of AMSPEC and MODIS at  
413 SOA and DF-49 (Figure 6, data from the Aqua and Terra spacecrafts combined). Figure 6A and B shows  
414 AMSPEC  $PRI_{12}$  observed at SOA compared to  $PRI_{12}$  sampled by MODIS using swath data(Figure 6A) and  
415 the gridded reflectance product (Figure 6B). Only little difference was found in the strength of the  
416 regression of these two datasets ( $r^2=0.63$  and  $r^2=0.60$  for swath and gridded data, respectively;  $p<0.01$ ).

417 Highly significant relationships between AMSPEC PRI<sub>12</sub> and MODIS PRI<sub>12</sub> were also found for the DF-49  
418 site. As with the Old Aspen site, only little differences were observed in the significance of the regression  
419 when using swath (Figure 6C) and gridded reflectance data (Figure 6D) ( $r^2=0.54$  and  $r^2=0.51$  for swath  
420 and gridded data, respectively;  $p<0.01$ ). However, MODIS observations sampled at  $\theta_v > 45^\circ$  were  
421 excluded from this dataset as previous research (Hilker et al., 2008b) has shown that owing to the  
422 increased pixel size at larger off-nadir angles, MODIS PRI<sub>12</sub> will be confounded by observations of clear-  
423 cuts and other non-forested elements.

424 Figure 7 shows the LiDAR derived CHM observed at SOA (Figure 7A) and DF-49 (Figure 7B). The extent of  
425 each raster approximates the largest possible viewing area of AMSPEC at each site ( $\theta_v = 78^\circ$ ). The  
426 respective towers are located in the center of each CHM. Notable differences can be observed in the  
427 structure of the canopy surface shown in Figure 7A and B with likely implications for  $\alpha_s$  estimated at DF-  
428 49 and SOA. The colors illustrate the differences in height, the larger gaps visible in Figure 7A are due to  
429 bogs found at the SOA site. Figure 8 gives an example of a hillshade model used to determine  $\alpha_s$  as a  
430 function of the solar position. Figure 8A represents a hillshade modelled from observations made at the  
431 Old Aspen site; Figure B shows the corresponding model for DF-49. Areas invisible to AMSPEC (=n.v.)  
432 have been eliminated from both hillshade raster by means of the viewshed algorithm. The solar  
433 geometry is identical in both examples ( $\theta_s = 45^\circ$ ,  $\phi_s = 170^\circ$ ), the approximate instantaneous field of  
434 view of one AMSPEC observation ( $\theta_v = 78^\circ$ ,  $\phi_v = 225^\circ$ ) has been illustrated as a superimposed ellipse  
435 (Figure 8A and B). While the majority of the canopy was visible at shorter ranges from the tower (<30m),  
436 the lower elements of more distant canopy surface areas were increasingly hidden behind other canopy  
437 parts and therefore no longer visible AMSPEC. The decrease in visible canopy area with distance from  
438 the tower area was rapid especially at the DF-49 site where the triangular crown shape allowed a view  
439 only of the tree tops at greater distances from the tower. Figure 9 shows the relationship between  
440 AMSPEC observed PRI reflectance and LiDAR estimated  $\alpha_s$  during one radiometer sweep (15 minutes)

441 observed under clear sky conditions. Both sites showed a strong linear correlation between PRI and  $\alpha_s$ ,  
442 which was however, more significant at DF-49 (Figure 9B) than at SOA (Figure 9A). The range of  $\alpha_s$   
443 between reflectance hotspot and darkspot was about three times bigger at DF-49 than at the SOA site.  
444 Both examples chosen in Figure 9 were sampled under similar physiological conditions ( $\varepsilon = 0.45 \text{ gCMJ}^{-1}$ ),  
445 roughly an hour before solar noon. Figure 10 shows  $\Delta \text{PRI} \Delta \alpha_s^{-1}$  as a function of EC-measured  $\varepsilon$  acquired  
446 during clear days at DF-49 and SOA, respectively. At both sites, a strong logarithmic relationship was  
447 found between  $\Delta \text{PRI} \Delta \alpha_s^{-1}$  and  $\varepsilon$ . The solid line shows the regression between  $\Delta \text{PRI} \Delta \alpha_s^{-1}$  and  $\varepsilon$  at the DF-  
448 49 site, the dashed line corresponds to the data acquired at SOA. The gray areas correspond to the 95%  
449 confidence interval around both regressions. Both regressions are falling within the 95% confidence  
450 interval of each other.

451

#### 452 **4. Discussion**

453 This study compared stand and satellite-scale assessments of PRI and  $\text{PRI}_{12}$  across two climatically and  
454 structurally different forested biomes. The webcam-based approach of Richardson et al (2007) was  
455 successfully used to quantify plant phenology and allowed a more objective selection of the study  
456 periods at DF-49 and SOA. While the focus of this study was on spatial scaling of PRI and  $\text{PRI}_{12}$ , a  
457 separate study will address potential seasonal changes in the  $\varepsilon$ :PRI relationship. For instance, the ratio  
458 of photosynthetic to non-photosynthetic material is expected to be important driver of canopy level  $\varepsilon$   
459 (Hall et al., 2008) especially at SOA, where C-uptake early in the year is expected to be driven largely by  
460 changes in springtime phenology and leaf green-up (Barr et al. 2004; Richardson et al. 2009). While only  
461 one camera position was used in this study to minimize the directional and background effects  
462 (Richardson et al. 2007), the multi-angular view of the webcam can potentially provide more  
463 information in this respect, as for instance the understory should be more visible from smaller zenith

464 angles, thus providing more prominent features in  $2G_{RB_i}$  value earlier in the year (Figure 2A,  
465 DOY<180).

466 At both sites, a strong, non-linear relationship existed between PRI and  $\varepsilon$  and PRI<sub>12</sub> and  $\varepsilon$  throughout the  
467 study period (Figure 3). This is an important finding as it demonstrates that the same method to  
468 separate physiological and directional effects in PRI is applicable across two structurally and climatically  
469 very different forest stands. This finding may also point towards a more generic application of this  
470 algorithm, at least in forested biomes, as numerous other studies have demonstrated the principal  
471 relationship between PRI and  $\varepsilon$  (Gamon et al. 1993; Gamon et al. 1997; Penuelas et al. 1997) at the leaf  
472 and stand-level scales.

473 While the PRI measurements under conditions where photosynthesis is not down-regulated (high  $\varepsilon$ ) is  
474 similar at SOA and DF-49 (Figure 3A and 3B), the Douglas-fir dominated stand exhibited lower PRI-values  
475 under situations where  $\varepsilon$  is low. Similarly, the PRI<sub>12</sub> measurements shown in Figure 3C and D are higher  
476 at the DF-49 site than at SOA when photosynthesis is less limited by  $\varepsilon$ . This is consistent with the lower  
477 amount of canopy shading observed at SOA (Figure 7,8) and also agrees with the results found in Figure  
478 9 and 10. The difference between sunlit and shaded PRI was more distinct at the DF-49 site than at SOA,  
479 which is consistent with the larger range in  $\alpha_s$  found at the coniferous site.

480 Figure 4 demonstrated the suitability of the Li-Sparse and Ross-Thick kernels to model the AMSPEC PRI<sub>12</sub>  
481 reflectance during a half hour interval at SOA and DF-49, thereby allowing a directional adjustment of  
482 the spectral observations to MODIS reflectance. The greater range of view zenith angles provided by  
483 AMSPEC II compared to the prototype version (Hilker et al. 2007) allowed a greater stability of the BRDF  
484 model with respect to predicting changes in PRI as a function  $\theta_v$ . This is critical especially when  
485 adjusting AMSPEC's geometry to that of satellite data, which, at least for high satellite elevations,  
486 cannot be accomplished through direct measurements alone. Also, direct comparisons of measurement

487 taken under identical viewing geometries is not necessarily desirable as 1) soil background reflectance  
488 effects may confound AMSPEC observations taken at small zenith angles and 2) a modelled reflectance  
489 based on several hundred observations obtained from different locations around the tower can provide  
490 a more realistic representation of the stand level reflectance, which is especially critical when scaling to  
491 moderate resolution sensors such as MODIS.

492 Only small differences were found in the significance of the relationship between PRI and  $\varepsilon$  and PRI<sub>12</sub>  
493 and  $\varepsilon$  at both sites. This is consistent with previous studies (Drolet et al. 2005; Drolet et al. 2008; Gamon  
494 et al. 1992; Hilker et al. 2009b; Middleton et al. 2009) and confirms the use of 551nm as a possible  
495 alternative to the commonly used reference wavelength at 570 nm. The comparison between MODIS  
496 and AMSPEC derived bands 11 and 12 presented in Figure 5 demonstrates the significant correlation  
497 between satellite data and BRDF corrected AMSPEC observations and also confirms the findings in  
498 Figure 4, which showed the suitability of the selected LSRT model to adjust directional differences  
499 between AMSPEC and MODIS reflectance during a half hour interval. The results shown in Figure 5 are  
500 also a rigorous assessment of the quality of MAIAC used to correct for atmospheric effects in MODIS  
501 band 11 and 12 (Lyapustin and Wang 2009; Lyapustin 2005) as they demonstrate that MAIAC allows a  
502 direct comparison not only of the normalized difference between two bands (Figure 6), but also of  
503 absolute reflectance. It should be noted, however, that there are differences in brightness observed by  
504 AMSPEC and MODIS at the SOA site (Figure 5A, 6A-B). One possible explanation could be variations in  
505 atmospheric conditions, as the BRDF measured by AMSPEC does also include diffuse illumination  
506 components, which will vary as function latitude because of differences in path length through the  
507 atmosphere.

508 The results shown in Figure 6 confirm previous findings from the DF-49 (Hilker et al. 2009b) and SOA site  
509 (Drolet et al. 2005; Nichol et al. 2000) and demonstrate that spaceborne assessments of  $\varepsilon$  are possible  
510 at least across these two biomes. The data shown in Figure 6 include forward and backscatter

511 observations from the Aqua and Terra spacecrafts combined. This is a significant advancement from the  
512 initial results found at SOA (Drolet et al. 2005; Drolet et al. 2008) and underlines the need for a careful  
513 consideration of atmospheric and directional impacts on PRI reflectance, which can confound the subtle  
514 changes in reflectance induced by physiological changes of the canopy. Almost no differences were  
515 found in strength of the correlation between AMSPEC and MODIS observed PRI<sub>12</sub> reflectance when  
516 considering gridded or swath data. This result was expected for SOA as this stand is quite large and  
517 homogeneous due to its location inside Prince Albert National Park. As a result, not many changes are to  
518 be expected in the neighbouring pixels around the tower. From the experiences of earlier studies (Hilker  
519 et al. 2009b) MODIS observations sampled at  $\theta_v > 45^\circ$  were excluded from the analysis of DF-49 data,  
520 which effectively reduced also the origin of the pixels to a smaller area around the tower thereby  
521 minimizing the effects of surrounding harvesting activities and clearcuts.

522 The results shown in Figure 7 and 8 demonstrate the notable differences in canopy shading observed at  
523 SOA and DF-49. The hillshade model used in this study was a simple, yet effective proxy of the daily and  
524 seasonal cycles in canopy illumination (Hilker et al. 2008b) (Figure 9). It should be noted that the  
525 hillshade approach only accounts for mutual shading effects and is therefore only an approximation of  
526 the radiation regime at a given time (Hilker et al. 2008b). Additionally, the model does not account for  
527 diffuse radiation conditions, and as a result, can only be used under clear sky conditions (One possible  
528 approach to extend this method for observations made under cloudy conditions would be to weight the  
529 model by the proportion of direct to diffuse irradiance, this is, however, of less interest when validating  
530 spaceborne observations). Consequently, this technique should not be considered as an absolute  
531 measure of canopy shading, it does, however, previous results have confirmed that it still yields realistic  
532 observations of the relative change in  $\alpha_s$  (Hall et al. 2008; Hilker et al. 2008b). The results shown in  
533 Figure 9A and B are consistent to those shown in Hall et al (2008) and demonstrate the dependency of  
534 PRI on  $\alpha_s$  during one radiometer sweep at DF-49 and SOA. While significant relationships existed at both

535 study areas, the correlation was stronger at the DF-49 site, which is consistent with the fact that canopy  
536 shading is much more predominant in the coniferous than at the deciduous stand. The impact of canopy  
537 shading on the stand level radiation regime can also be observed when comparing the range of shadow  
538 fractions during one radiometer sweep at SOA and DF-49 (Figure 9).

539

## 540 **5. Conclusions**

541 The slope of the relationship between  $\alpha_s$  and PRI ( $\Delta\alpha_s\Delta\text{PRI}^{-1}$ ) is a very similar logarithmic function of  $\epsilon$   
542 (Figure 10) for both sites. The parameters of the two functions do not differ significantly, suggesting  
543 that one function can describe two very different vegetation communities, in two very different  
544 climates. This is a key finding of this study. First, it confirms that the changes in PRI reflectance at SOA  
545 and DF-49 were both driven by physiological changes in the canopy rather than extraneous effects, as  
546 demonstrated in the inference framework introduced in Hall et al. (2008). Secondly, it can be concluded  
547 from Figure 10 that when viewing the canopy at one angle, as is the case with MODIS, the differences  
548 observed in the relationship between PRI and  $\epsilon$  at DF-49 and SOA (Figure 4) can be attributed mainly to  
549 differences in the canopy structure and shadow fraction. This finding is consistent also with previous  
550 studies (Barton and North 2001; Sims and Gamon 2002) and emphasizes the effect of canopy structure  
551 on PRI (Hall et al. 2008; Middleton et al. 2009). It shows that single date remote sensing of  $\epsilon$  at a single  
552 view angle will need to take into account the ratio of photosynthetically active to non-photosynthetic  
553 canopy elements, and shadow fraction.

554 Importantly, the study shows that instantaneous spectral measurements of a canopy at multiple view  
555 angles, which are possible using a sensor viewing the canopy along track, as the Chris sensor aboard the  
556 Proba platform, could measure both  $\alpha_s$  (using visible and NIR bands with mixture decomposition as in  
557 Hall et al. 1995) and PRI (using the 531 and 570 nm bands for the different view angles). Along any  
558 orbital track an instantaneous estimate of  $\Delta\alpha_s\Delta\text{PRI}^{-1}$  could then be computed for each pixel in the scene,



559 hence canopy  $\epsilon$  could be inferred with a functionally invariant logarithmic relationship across divergent  
560 biomes. Adding the NDVI bands to such a sensor to measure  $f_{PAR}$  could provide a direct estimate of GPP.  
561 In the same way our results show that use of an AMSPEC like instrument suite atop a tower, can directly  
562 measure LUE,  $f_{PAR}$  and GPP as an adjunct to eddy-correlation measures of NEE and NPP. The advantage  
563 of the AMSPEC approach is that it measures GPP directly without the need for measuring respiration. As  
564 a result, differencing AMSPEC measures of GPP and eddy-correlation measures of NPP could provide an  
565 independent means for inferring respiration without resorting to measurements of night time fluxes  
566 (Jassal et al. 2007).

567 We therefore propose a field campaign including multiple AMSPEC- like instruments to compare  
568 continuous PRI measurements and EC-flux data thereby helping to calibrate coarser scale observations  
569 to tower-based measurements and assessing the potential for a generic model of PRI across different  
570 vegetation and land-cover types.

571

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580

581

582 **Figures**

583 Figure 1: In-situ photograph of AMSPEC II taken at the Old Aspen site. The system features a pan-tilt unit  
584 which allows the sensor head to be moved at any zenith angle between 40 and 78° at a view azimuth  
585 between 0 and 360°. The upward looking sensor features a cosine diffuser to correct for varying solar  
586 altitudes. Canopy reflectance is determined from solar irradiance and canopy radiance. A webcam  
587 picture is automatically taken with every spectrum that is sampled. An identical system has been  
588 installed at the DF-49 site.

589  
590 Figure 2A-B: Estimate of the canopy phenology as observed from AMSPEC's webcam (Richardson et al.,  
591 2007). Figure 2A: Spring green-up and leaf down of the Old Aspen site as estimated from 2G\_RBi  
592 ( $\theta_v=73^\circ$ ,  $\phi_i=65^\circ$ ). Green-up and leaf-down were quantified using null and inflection points of the  
593 polynomial fit. The missing data is due to a downtime of AMSPEC at DF-49 between DOY 197 and 231.

594  
595 Figure 3A-B: Relationship between AMSPEC-observed PRI and EC-measured  $\epsilon$  for hotspot and darkspot  
596 reflectance (sunlit and shaded components of the canopy, averaged to daily observations). The SOA site  
597 is shown in Figure 3A, Figure 3B represents the PRI-  $\epsilon$  relationship at DF-49. Figure 3 C-D: Relationship  
598 between AMSPEC -observed  $PRI_{12}$  and EC-measured  $\epsilon$  at SOA (3C) and DF-49 (3D).

599  
600 Figure 4A-B: Example of a BRDF model during one MODIS overpass (spacecraft noon  $\pm$  15 minutes).  
601 Figure 4A shows data from SOA Figure 4B is based on observations taken at DF-49. The x and y-axis  
602 represent the planar coordinates (origin=tower), the z axis represents the  $pPRI_{12}$  value at this location.

603 The black dots represent the actual  $\rho\text{PRI}_{12}$  measurements observed by AMSPEC, the black lines show the  
604 residuals to the fitted surface. The red dots are the PRI observations taken by AMSPEC while tracking the  
605 flight path of EO-TERRA. The yellow dot (blue dot in Figure B) represents the corresponding zenith and  
606 azimuth for the actual MODIS observation. The green dots in Figure 4A represent AMSPEC observations  
607 taken with  $\theta_z = \theta_v$ . No solar tracking was done during this overpass at DF-49 as  $\theta_i$  exceeded the range of  
608 observable  $\theta_z$ .

609

610 Figure 5A-B: Comparison of AMSPEC observed, BRDF adjusted MODIS bands 11 and 12 and MODIS  
611 observed reflectance at Band 11 and 12 (2009 data, Aqua and Terra combined). Figure 5A shows  
612 observations taken at SOA, Figure 5B shows the reflectance observed at DF-49. The second y-axis for  
613 MODIS Band 12 was introduced for illustration purposes.

614

615 Figure 6A-D. Comparison between AMSPEC observed, BRDF adjusted  $\text{PRI}_{12}$  and MODIS  $\text{PRI}_{12}$ . Figure 6A  
616 and B show the results for non-gridded (swath) data and gridded data, observed at SOA Figure 6C and D  
617 show the corresponding results for the DF-49 site

618

619 Figure 7A-B: LiDAR derived canopy surface model (CSM) observed at SOA (A) and DF-49 (B). The spatial  
620 extend of the models approximates the maximum viewing area of AMSPEC ( $\pm 150$  m from center=tower).  
621 The line with higher elevations shown Figure 7A is due to a tram line which was installed at the site  
622 during the BOREAS field experiment. This area has been excluded from the analysis of  $\alpha_s$ .

623

624 Figure 8: LiDAR derived viewshed model for  $\theta_v=45^\circ$  as observed at the SOA (Figure A) and DF-49 site  
625 (Figure B). The hillshade analysis was carried out only for those canopy areas visible from the tower. The  
626 ellipse shown in both figures represents an example of an area observed AMSPEC II at a given zenith and  
627 azimuth (here:  $\theta_v=78^\circ$  and  $\phi_v=225^\circ$ ). The relatively smooth canopy at the SOA site yields an almost  
628 complete observation of the canopy around the tower with  $\alpha_s$  being relatively small (here  $<30\%$ ), where  
629 as  $\alpha_s$  is much higher at DF-49.

630

631 Figure 9: Relationship between AMSPEC observed PRI and  $\alpha_s$  observed during one 15-minute interval  
632 (Figure A: SOA, Figure B: DF-49). During this time period  $\varepsilon$  was assumed to be constant ( $\varepsilon=0.45 \text{ g CMJ}^{-1}$   
633 in both cases).

634

635 Figure 10: Relationship between  $\Delta \text{PRI}$   $\Delta \alpha_s - 1$  and EC-measured  $\varepsilon$ . The regression line established from  
636 DF49 data is solid; the one established from SOA data is dashed. The gray areas correspond to the 95%  
637 confidence interval around both regressions. Both regressions are falling within the 95% confidence  
638 interval of each other and both show a similar, logarithmic behaviour.

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# Upward looking probe with cosine diffuser

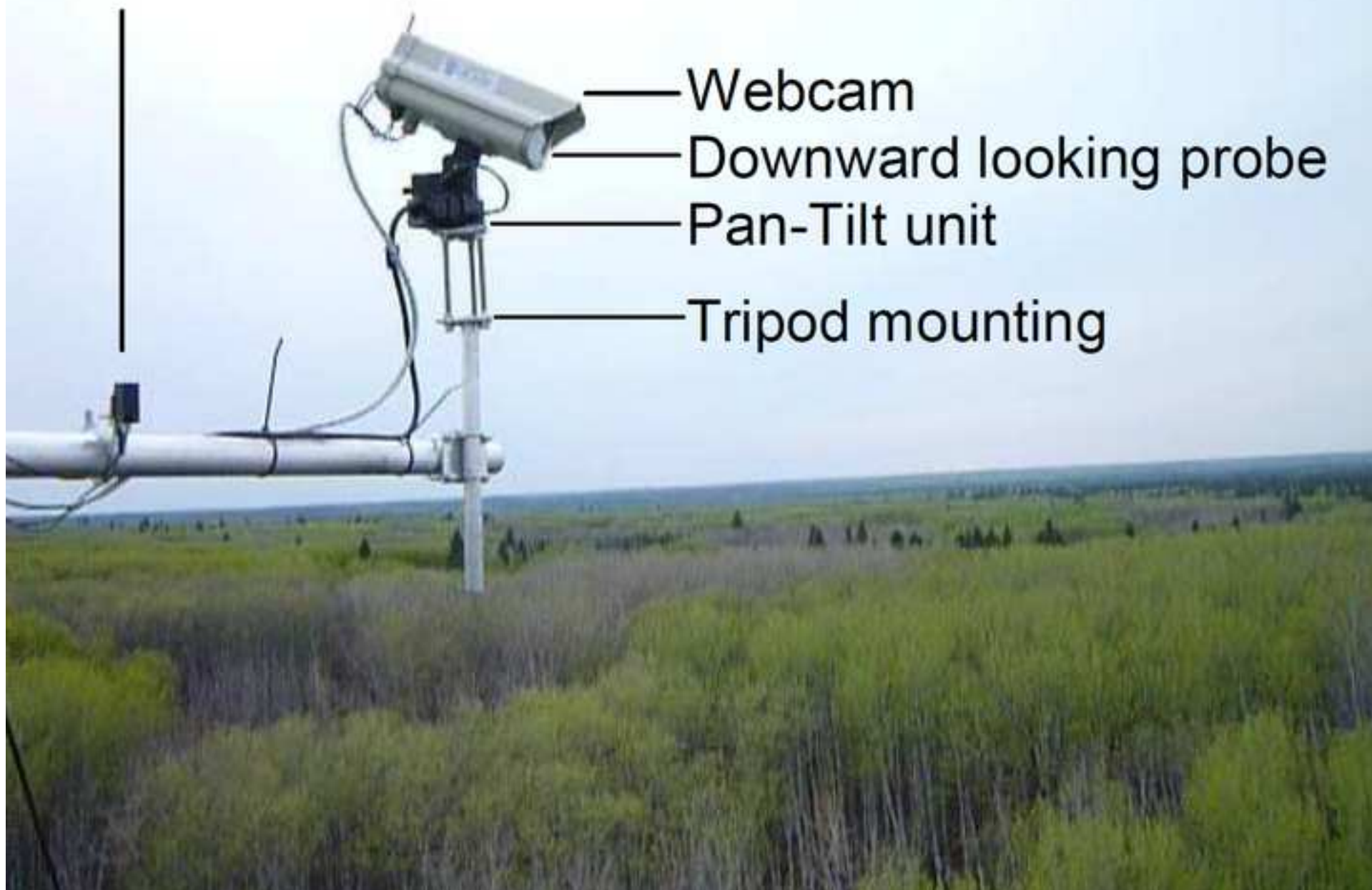


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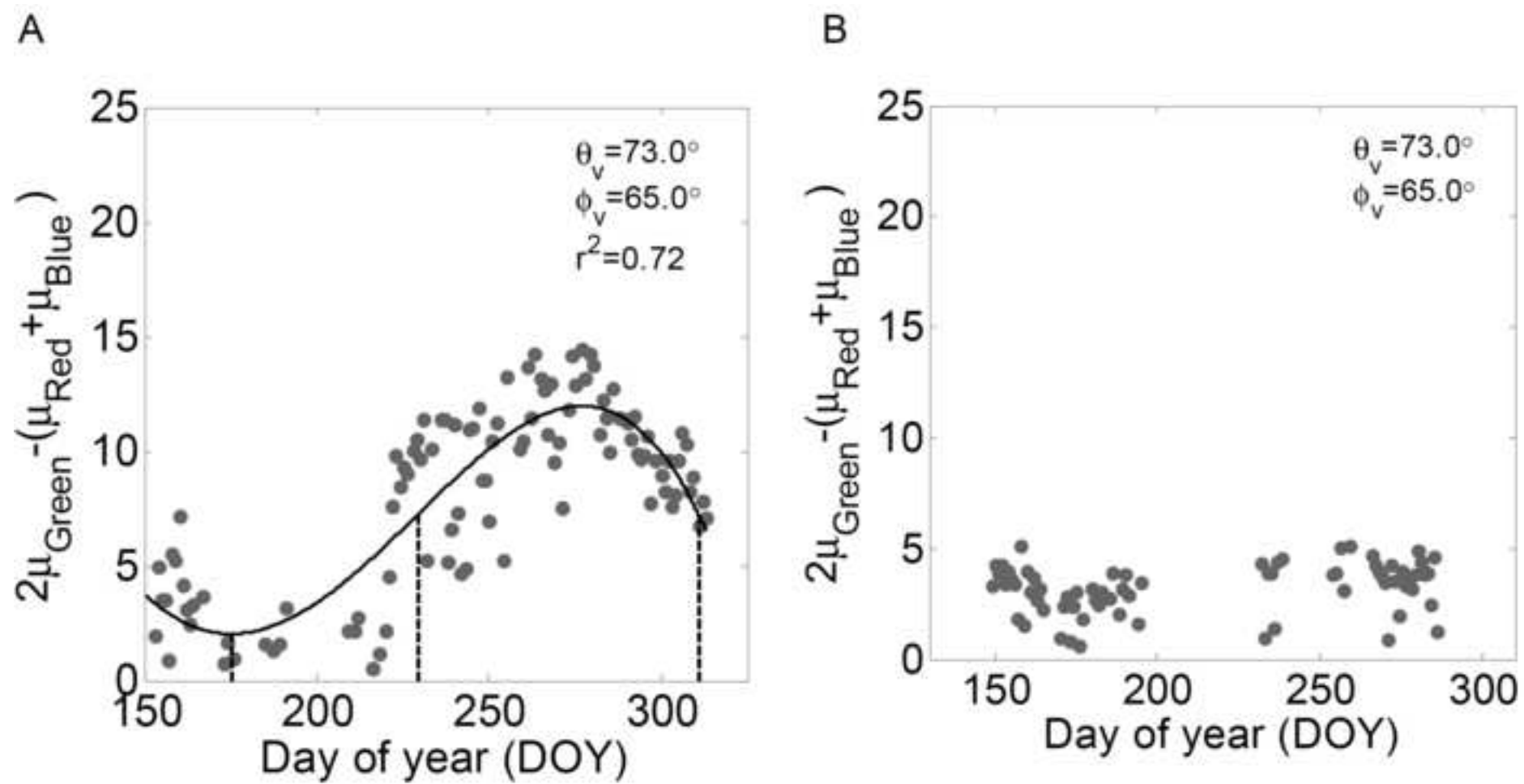


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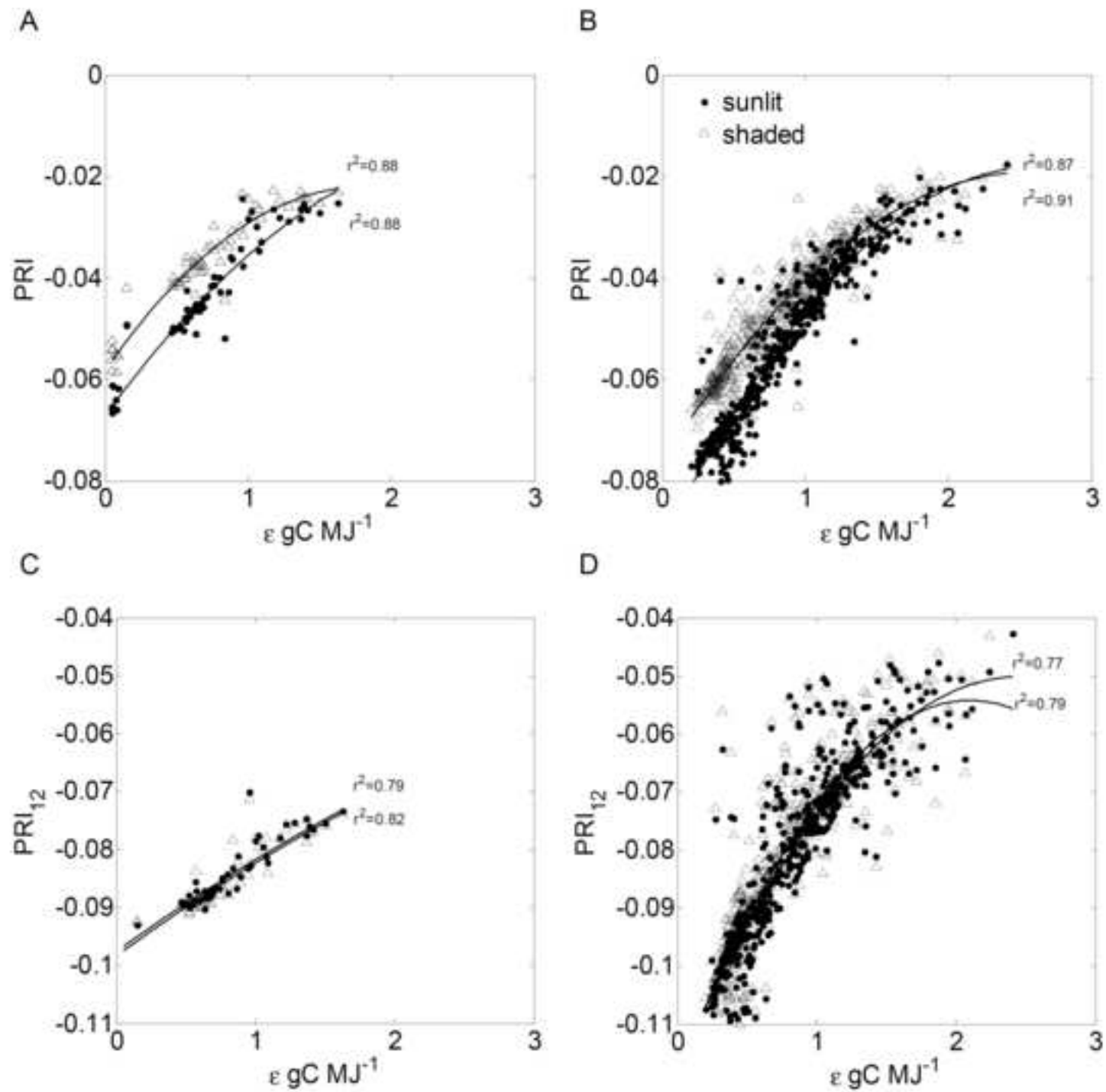


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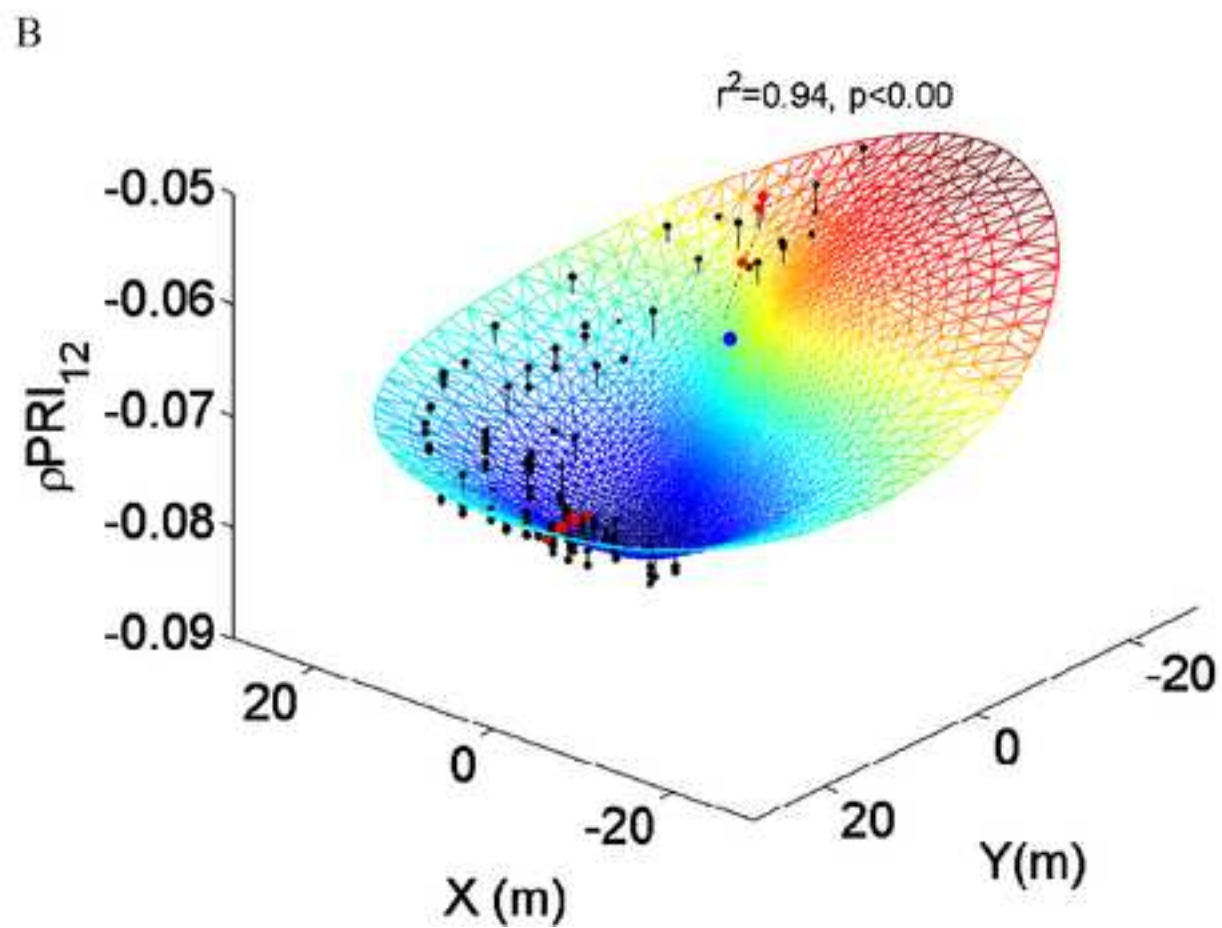
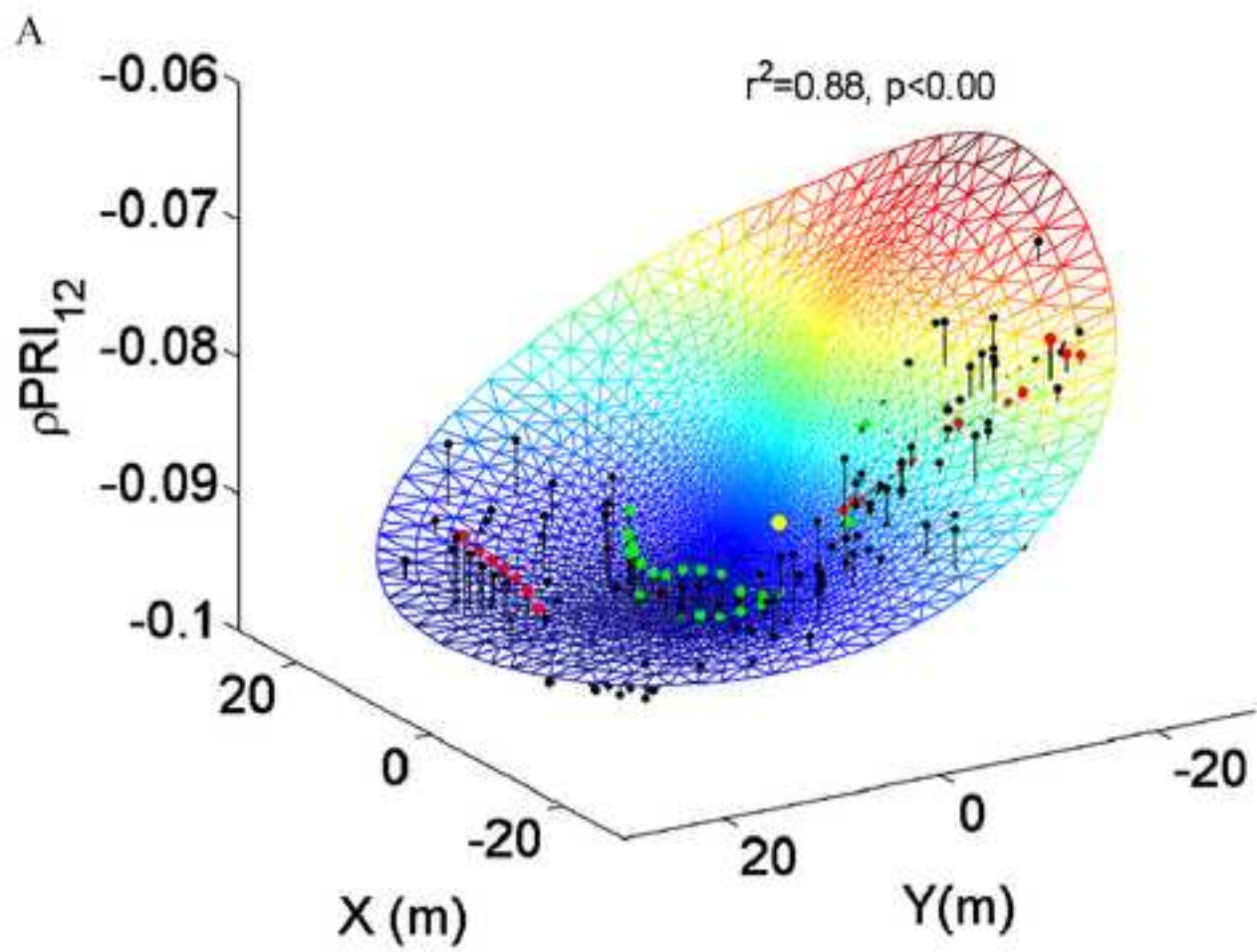


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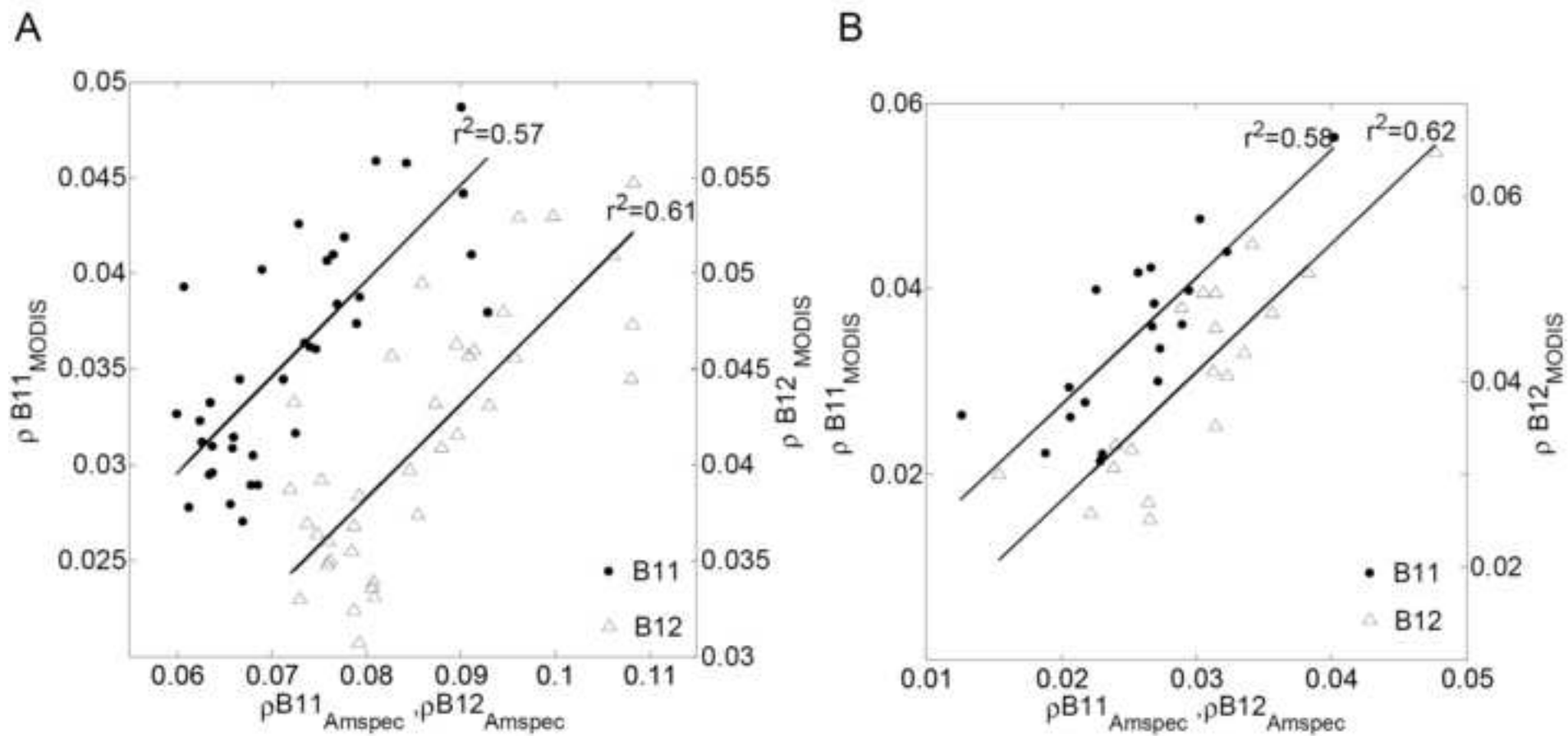


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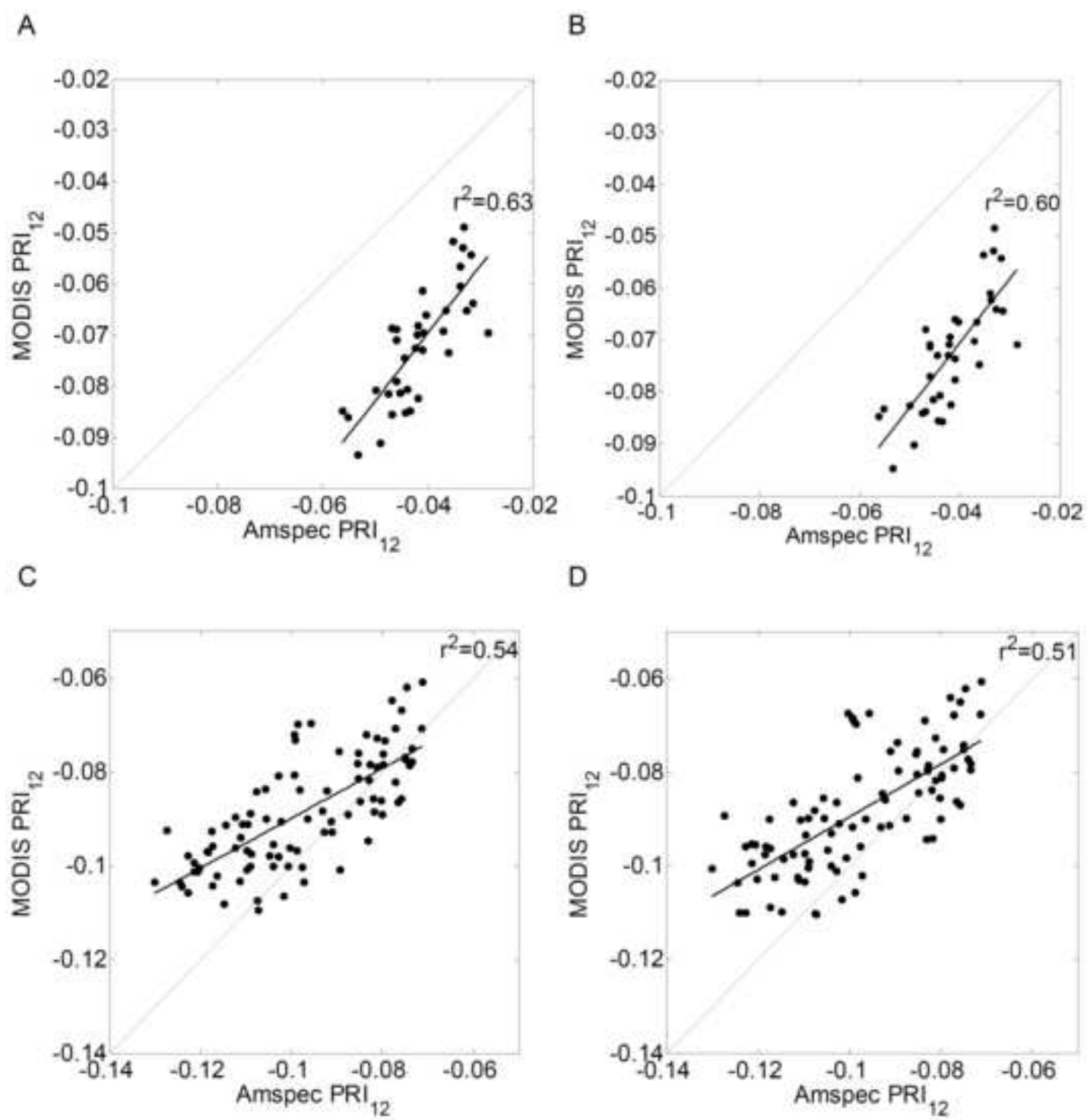
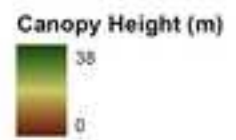
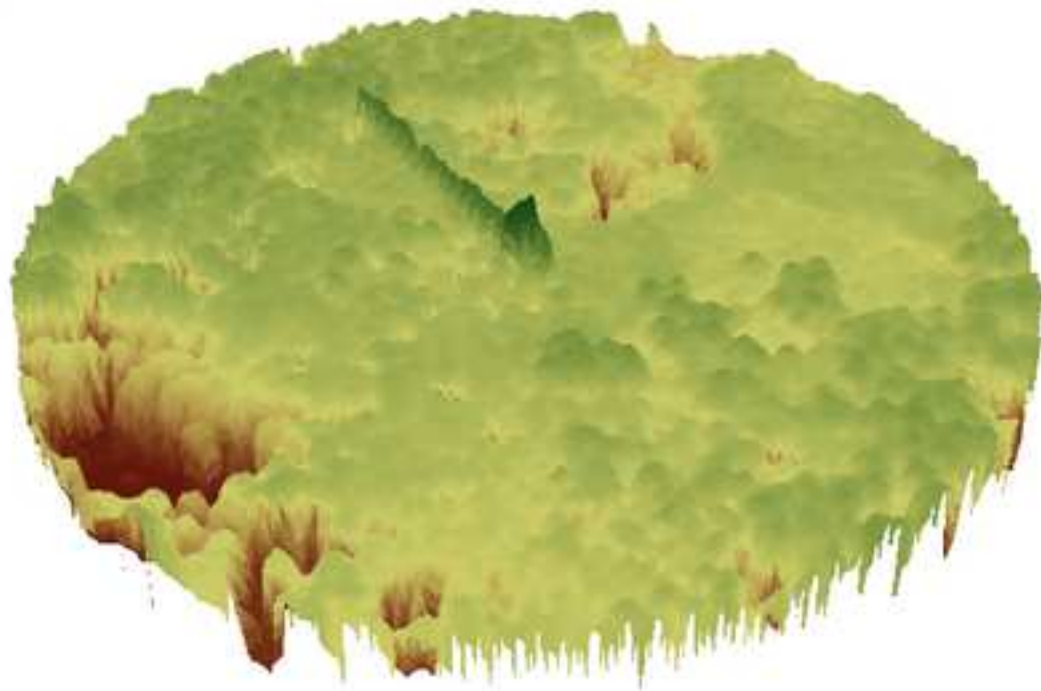


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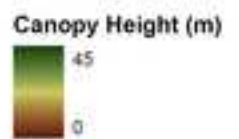
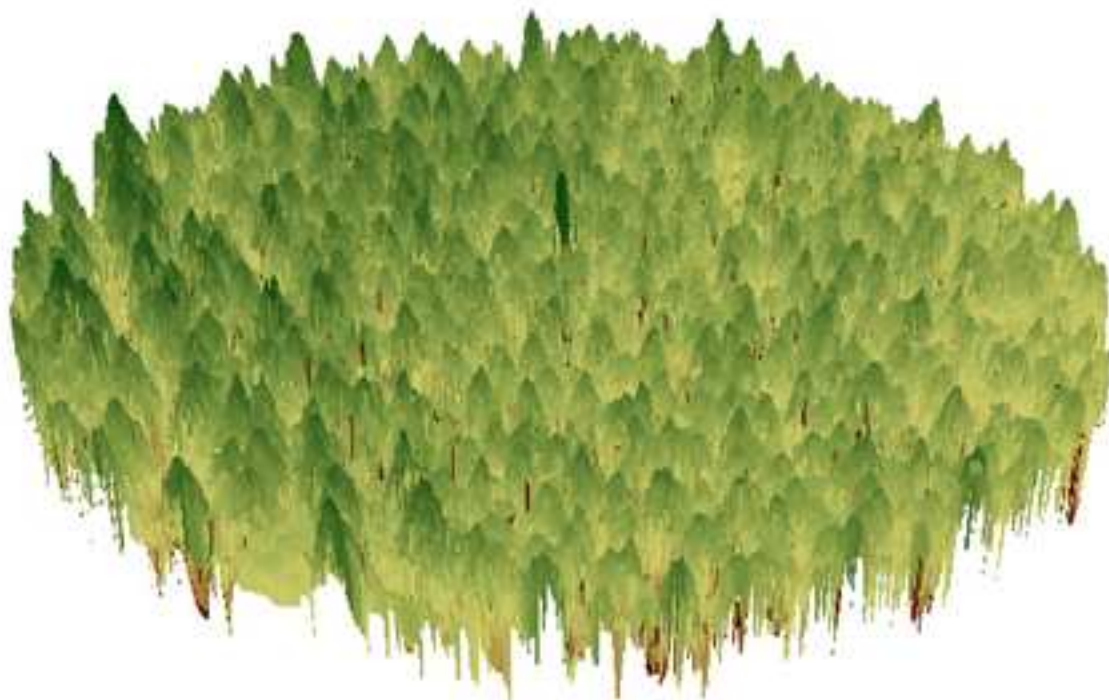


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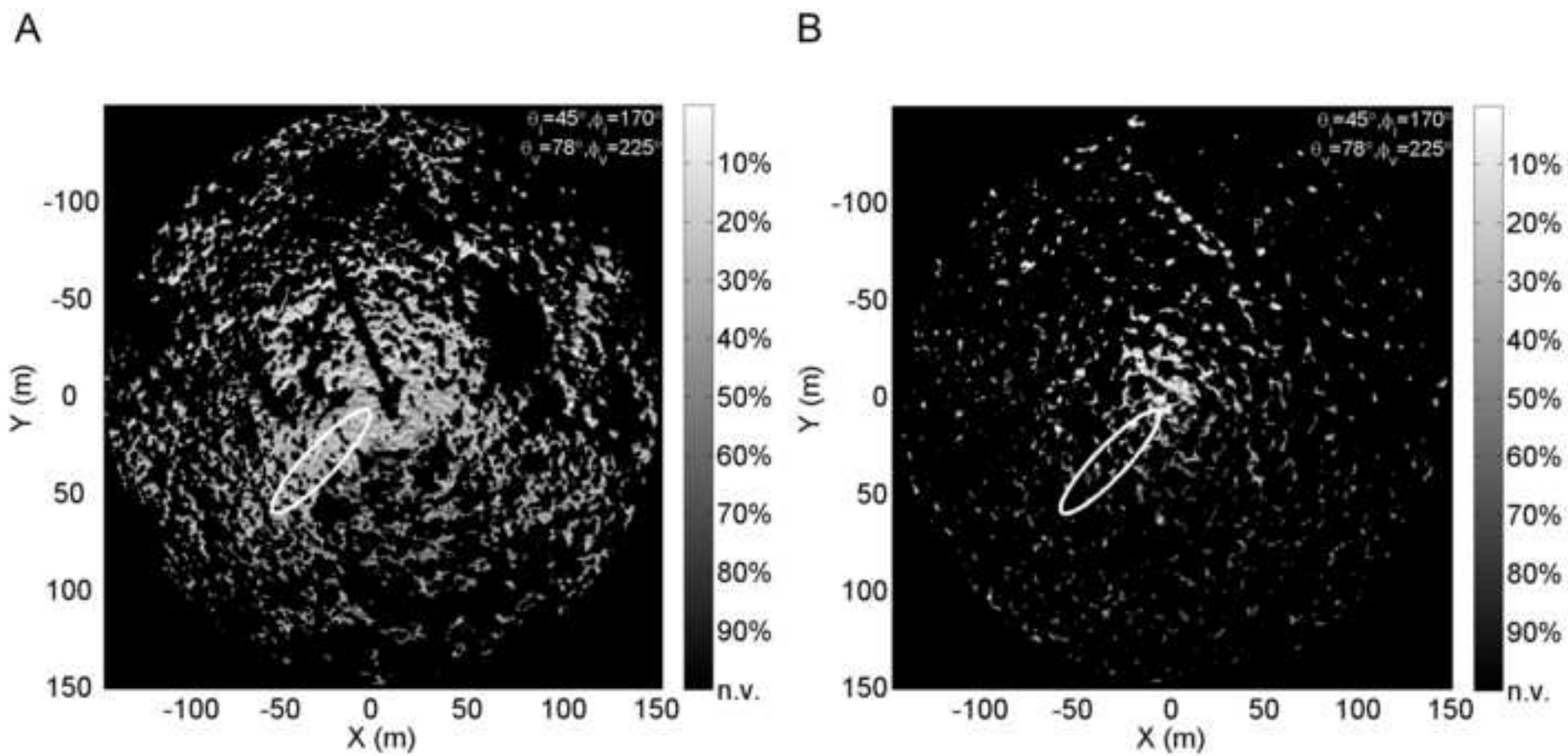




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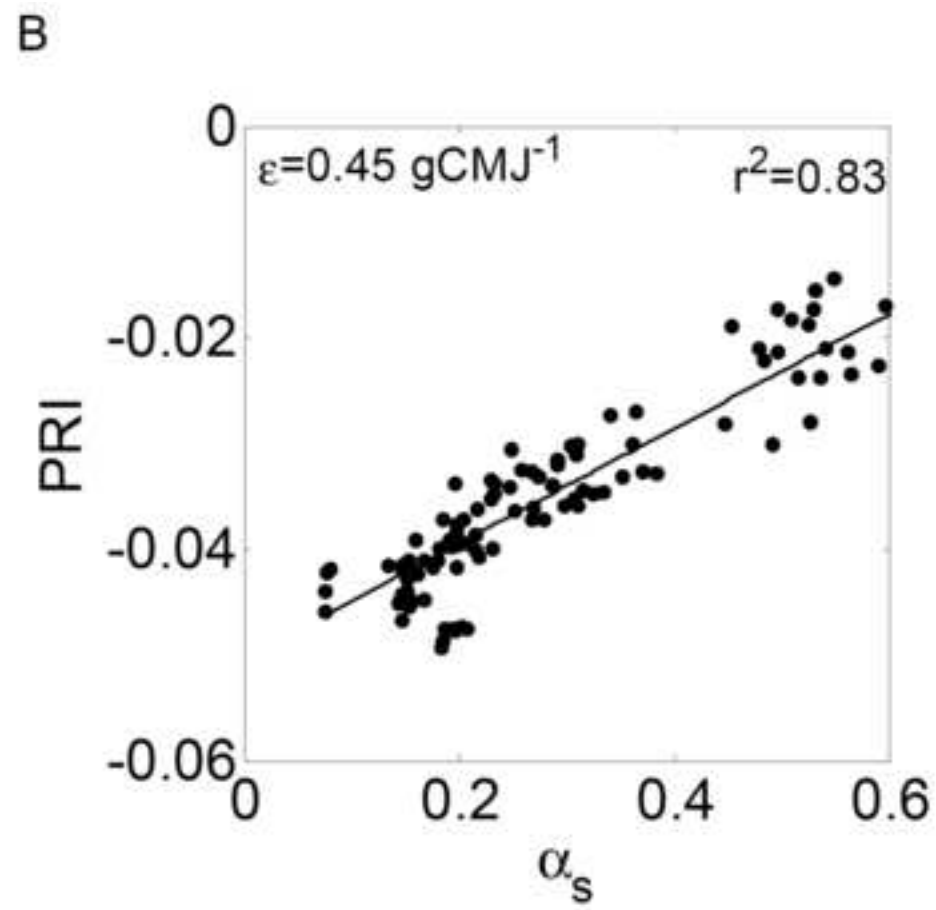
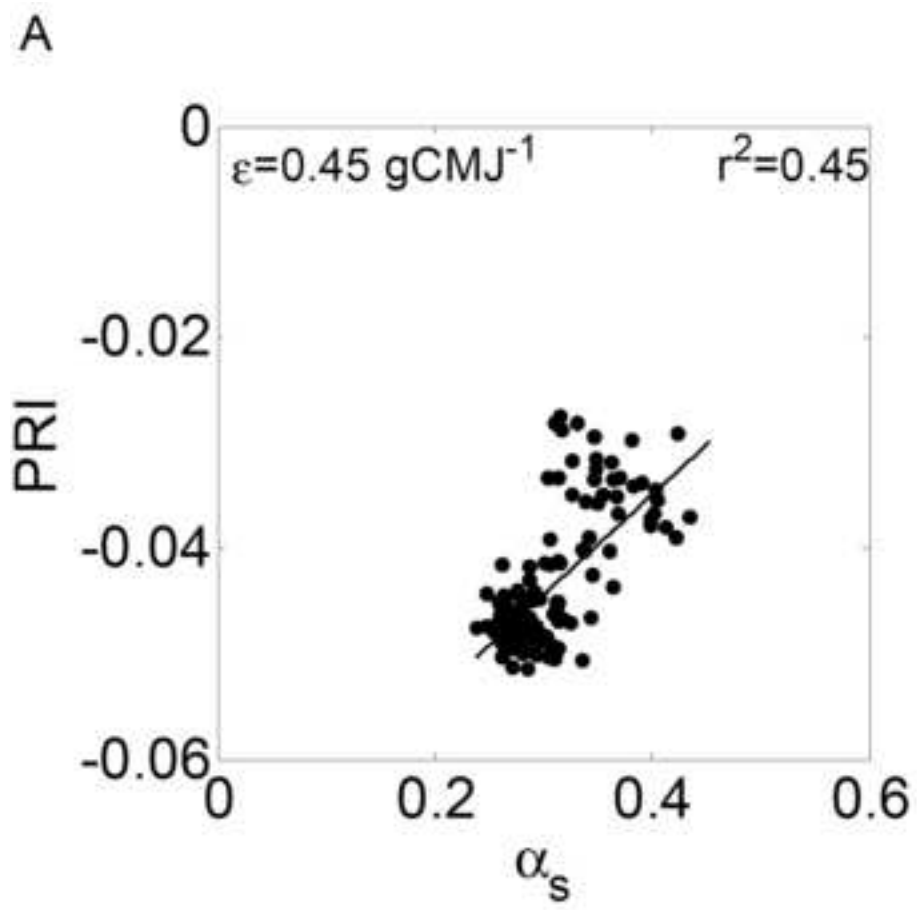


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